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Late Quaternary bedrock incision in the Narmada river at Dardi Falls

Avijit Gupta¹, Vishwas S. Kale^{2,*},
Lewis A. Owen³ and A. K. Singhvi⁴

¹School of Geography, University of Leeds, Leeds, LS2 9JT, UK

²Department of Geography, University of Pune, Pune 411 007, India

³Department of Geology, University of Cincinnati, Cincinnati, OH 45221, USA

⁴Planetary and Geosciences Division, Physical Research Laboratory, Ahmedabad 380 009, India

Fluvial incision in bedrock is common in many rivers of the Indian Peninsula. We investigated a site in the gorge of the Narmada river at Dardi Falls that displays geomorphic evidence of intense bedrock erosion. We

report here a terrestrial cosmogenic radionuclide date from an eroded rock surface in Peninsular India. Terrestrial cosmogenic radionuclide dating of the rock surface adjacent to the inner gorge indicated that the minimum age of the gorge is 40 ka. We suggest that the present gorge has developed in two phases, separated by a period of large-scale aggradation that filled the gorge with alluvium. Gorge formation is most likely associated with tectonic activity in the Son–Narmada–Tapi lineament zone. Erosion at this scale also requires large palaeodischarges with high unit stream power. This study illustrates the powers of combining newly developing geomorphic, DEM and geochronological methods to elucidate the dynamics and nature of landscape evolution.

Keywords: Bedrock erosion, cosmogenic radionuclide dating, DEM, gorge, Narmada.

THE erosion-dominant landscape of the Indian Peninsula is characterized by rivers downcutting into underlying bedrock. Many large rivers of Peninsular India flow through scablands, rock gorges and knickpoint-related waterfalls. However, the dates of such downcutting or rates of erosion are not known for any of the Peninsular rivers, with the exception of the pre-Holocene Tapi Gorge¹ and the historical evidence of rapid bedrock erosion by the Indrayani river². We report here a ^{10}Be terrestrial cosmogenic radionuclide surface exposure age determination for the Narmada Gorge at Dardi Falls and discuss its significance in terms of regional geomorphology and bedrock erosion.

The 1300-km long Narmada river in Central India flows through three major bedrock gorges separated by alluvial basins. In the downstream direction these are the Marble Canyon near Jabalpur, Punasa Gorge near Khandwa, and Dhadgaon Gorge west of Badwani³. The present study is on the Punasa Gorge, where at Dardi (22°18'49"N and 76°21'11"E) the Narmada river (Figure 1) has eroded a >100 m deep and 1–7 km wide gorge with an inner canyon⁴ (Figure 2). A ca. 10-m high waterfall occurs at the head of the inner canyon. Thick (>50 m) alluvial deposits of late Quaternary age overlie the bedrock on the right bank (Figure 2a), whereas the left bank rises to a set of strath terraces with abraded and water-polished rock surfaces. The back of the terraces is covered by riverine alluvium and soil. Similar rock surfaces are likely to extend below the alluvium that occurs on the top of the right bank.

Quartzites of the Vindhyan Supergroup constitute the local bedrock. To the south, a major fault runs parallel to the Narmada river⁵ (Figure 1c). The river flows through the large tectonic feature of the Son–Narmada–Tapi (SONATA) lineament zone characterized by neotectonism, moderate seismicity^{5–7}, and several longitudinal fault-bound blocks with an episodic history of vertical and lateral movements⁶. The resulting river morphology demonstrates tectonic and structural control as illustrated in Figure 1c and d, which is a digital elevation model (DEM) of the Dardi Falls area

*For correspondence. (e-mail: vskale@unipune.ernet.in)

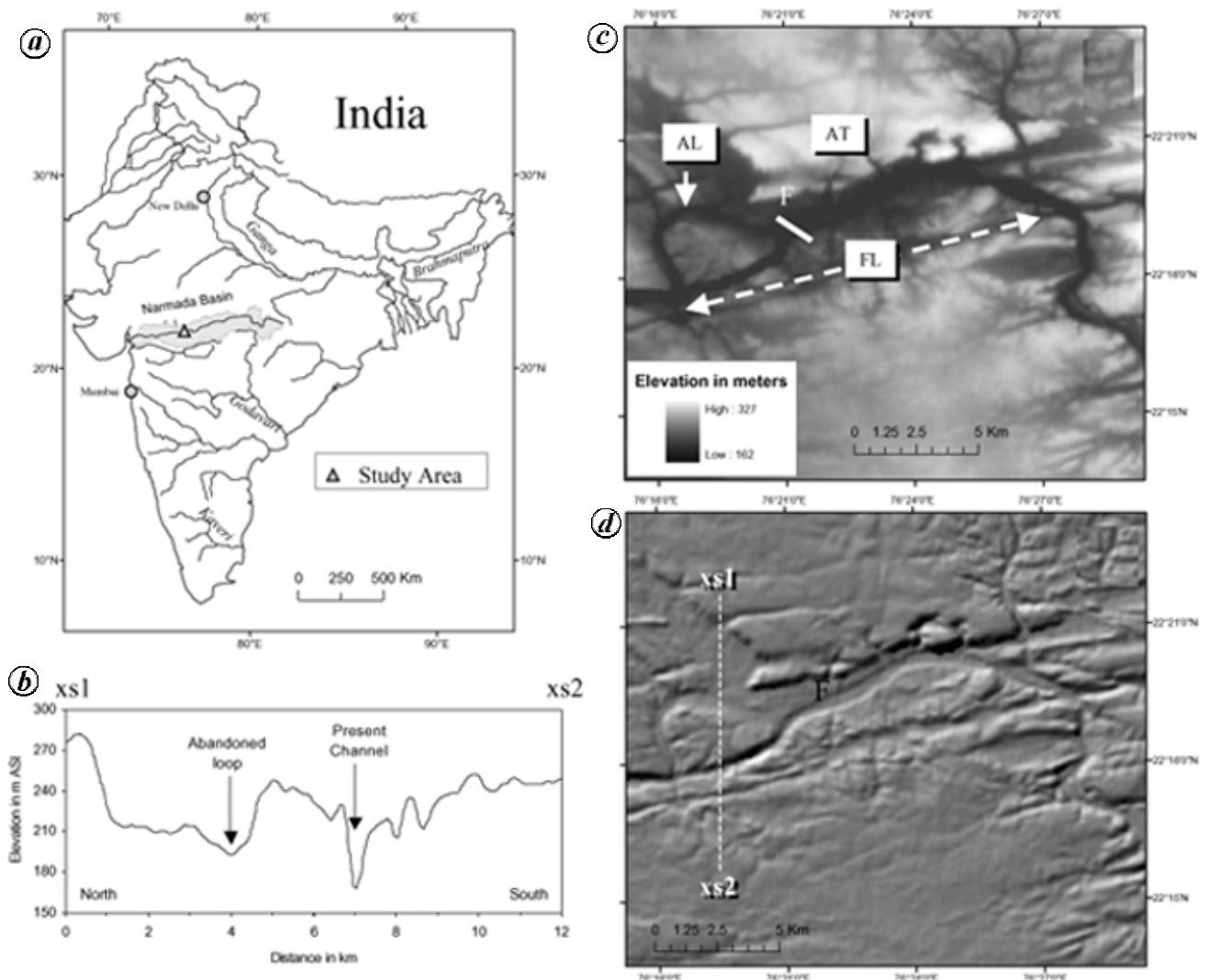


Figure 1. *a*, Location of the study area. *b*, Cross-profile across the abandoned loop and the modern channel of the Narmada river. *c*, Digital elevation model. *d*, Shaded relief map of the Dardi Falls area based on 90-m SRTM data. AL, Abandoned Loop; AT, Alluvial Terrace; F, Dardi Falls and FL, Major south fault. Location of the cross-profile is shown in *d*.



Figure 2. *a*, A view of the channel scabland area around Dardi Falls, the inner channel, and thick alluvial deposits on the right bank (view upstream). *b*, Inner channel of the Narmada river.

that was generated using 90 m Shuttle Radar Topography Mission (SRTM) data. An abandoned channel loop is present immediately downstream of the falls (Figure 1 *c* and *d*). The elevation of the abandoned channel is about 25 m higher than that of the present bed of the Narmada (Figure 1 *b*). This association suggests that the abandonment of the river loop, creation of the waterfall, and the inner gorge eroded by upstream retreat of the falls are interlinked. DEM analysis reveals yet another abandoned channel bend in the rock, ca. 9 km downstream of the Dardi Falls.

Samples were collected and processed for ^{10}Be terrestrial cosmogenic radionuclide surface exposure dating from the flat rock terrace surfaces away from the vertical walls of the gorge. Two sites on the left bank were selected for sample collection, one on the flat rock surface close to the low water level of the inner gorge (DF1) and another on the shoulder of the inner gorge (DF3). At each site an *in situ* river-polished slab of quartzite, ca. 5 cm thick, was collected from the surface with hammer and chisel.

Following crushing and sieving, quartz was separated from the 250–500 μm size fraction using the method of Kohl and Nishiizumi⁸. The hydroxides were oxidized by ignition in quartz crucibles. BeO was then mixed with Nb metal prior to determination of the $^{10}\text{Be}/^9\text{Be}$ ratios by accelerator mass spectrometry at the Center for Accelerator Mass Spectrometry, Lawrence Livermore National Laboratory. Isotope ratios were compared to ICN ^{10}Be standards (K. Nishiizumi, pers. commun.) using a ^{10}Be half-life of 1.5×10^6 years. The measured isotope ratios were converted to ^{10}Be concentrations in quartz using the total Be in the samples and the sample weights. Radionuclide concentrations were then converted to zero-erosion exposure ages using the scaling method of Stone⁹, and were corrected for shielding and geomagnetic field variations as described by Farber *et al.*¹⁰. The sample DF1 could not be processed because quartz had too many impurities and requires resampling. However, DF3 provided reasonable concentrations. The age data for the ^{10}Be terrestrial cos-

mogenic date are shown in Table 1 and in the following paragraphs we discuss the possible implications of this age, so as to initiate more research on the timing of bedrock incision in the region.

The exposure age of DF3 sample is at least 40 ka. This integrates in it the shielding effects of later burial and exhumation. During gorge formation, the river obviously had adequate power to sculpt large-scale fluvial forms in quartzite. The sample was from the rock terrace above the inner channel allowing us to infer that the inner gorge is younger than the determined date as it is cut into the sampled bedrock.

Remnants of the alluvial valley-fill deposits on the right bank of the Narmada river (Figure 2 *a*) indicate large-scale aggradation. No local radiometric dates of these alluvial deposits are available, but numerous radiocarbon dates from the central Narmada Basin^{11,12} suggest that this alluvium is late Pleistocene in age, ca. 9–40 ka. This alluvium occurs along the Narmada and has been identified from stratigraphic evidence seen as the deposits of a mobile, meandering river that carried large quantities of sand⁴. At Dardi, this alluvium filled at least partially a valley in quartzitic sandstone that varies in width between 1 and 7 km and is 100 m deep. The alluvium was then eroded to expose the present Punasa Gorge.

At this stage, based on a single age, we were not able to reconstruct the timing or origin of the canyon-like inner channel of the Narmada. We can consider two possible scenarios (Figure 3). In the first scenario, the inner gorge could have been formed in the same downcutting event as the entire gorge. The gorge subsequently was filled by alluvium of the meandering palaeo-Narmada, and later exposed by the river in a second downcutting episode of post-Pleistocene age when most of the alluvial cover was removed. Alternatively, the Narmada downcut a wide valley at >40 ka ago, which was then filled by late Pleistocene alluvium followed by a Holocene excavation during which the river had enough power to remove the alluvium and excavate the inner gorge in the rock.

Tectonically steepened gradient of the river may explain part of the incision^{13,14}, but substantive amount of water needs to flow down the Narmada on a steepened gradient in order to downcut a wide and deep channel in quartzite. The standard stream power equation directly relates power with both slope and depth of a river. This implies a time of large floods. The present Narmada is capable of moving at least 2 m boulder slabs in large floods (Figure 2 *b*). Past floods must have been bigger, as active gorge formation does not happen at present.

Hypothesis 2 (Figure 3) suggests that the following events occurred in succession: (i) Principal gorge downcutting event probably during isotopic stage 3 (60–40 ka), a period more humid than the present, seen in several records in India^{15–19}. (ii) Sediment production, transport and infill by a meandering palaeo-Narmada. (iii) Removal of sediment and downcutting of the inner gorge.

Table 1. Terrestrial cosmogenic surface exposure date and age for sample DF3

Sample number	DF3
Latitude ($^{\circ}\text{N}$)	22.31
Longitude ($^{\circ}\text{E}$)	76.35
Altitude (m asl)	190
Thickness of sample (cm)	5
Quartz mass (g)	15.1965
^9Be mass (g)	0.9808
Conc. ^9Be (mg/g)	0.437
$^9\text{Be}/^{10}\text{Be}$ (10^{-14})	6.130 ± 1.184
^{10}Be concentrates atoms/g SiO_2 (10^4)	11.40 ± 2.20
Shielding factor (Skyline 60° for 180° azimuth)	0.604347
Age not corrected for variations in palaeomagnetic field (ka)	47.5 ± 9.3
Age corrected for variations in palaeomagnetic field (ka)	39.7 ± 9.7

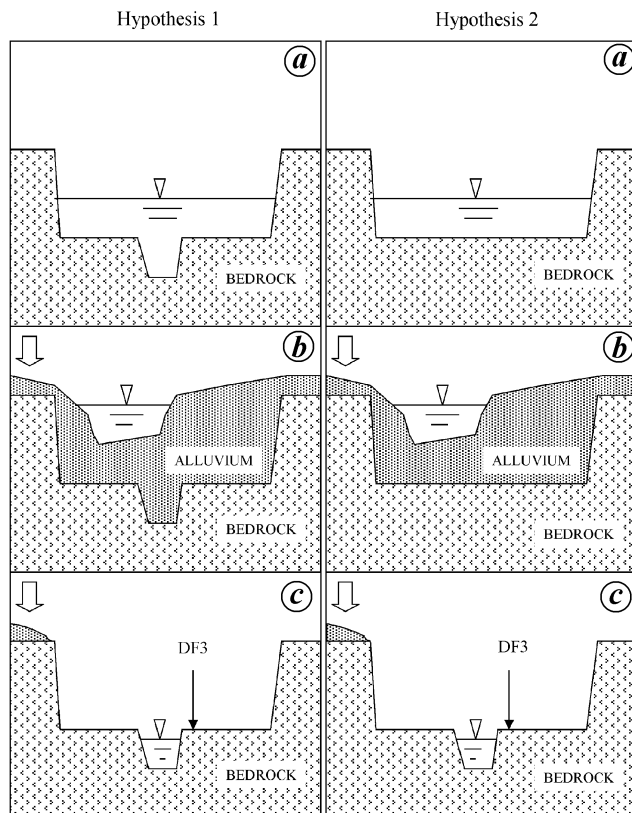


Figure 3. Diagrammatic sketch of two hypotheses of gorge formation at Dardi. See text for explanation. DF3, Dated sample.

The gorges of the Narmada apparently had complex history of tectonic movements and hydrological changes²⁰. Near Jabalpur, like the loops at Punasa, the river has abandoned a meander to flow into a straight 3-km gorge downstream of a waterfall in marble²¹. It is possible that in the past the river flowed across bends and excavated tectonics-related lineations in rock to form gorges. Abandonment of previous loops and course-straightening could be due to both tectonic triggers and large floods. Excavation of the gorge required deep floods with high velocity along weakened rock. Once confined in a gorge, a river does not change its course. We need more dates to properly reconstruct the evolution of the Narmada gorges, but it seems likely that besides tectonic control, gorges are the result of periods of flood-prone river hydrology. The intermediate channel-filling alluvium was deposited by a meandering river⁴ during a period of low floods and weak monsoon.

This is an initial study of the powers of combining newly developing geomorphic, DEM and geochronological techniques to help understand the dynamics and evolution of landscapes in the Indian Peninsula. It is hoped that this will stimulate similar future research in the region and regions outside India.

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